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短報

Abnormal Discharge of High-temperature Hot Spring Water Observed at the Abandoned Well in the Okuhida Hot Spring Area during the Earthquake Swarm Occurred at Gifu-Nagano District, Central Japan

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岐阜-長野地方の群発地震に伴って観測された 奥飛騨温泉郷の枯渇源泉の異常湧出

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要 旨

2020年4月下旬から長野県中部-岐阜県飛騨地方を震源とする地震が頻発し、この地震活動 に伴って奥飛騨温泉郷の枯渇源泉から高温泉が再湧出した.奥飛騨温泉郷は活火山の焼岳の西 方約4kmに位置するため、今回観測された高温泉の再湧出現象が焼岳の火山活動に起因する かどうかを調べるために、温泉水の主要化学組成および水素・酸素安定同位体組成を分析した. その結果、再湧出した温泉水は天水を主体とする中性のNa-HCO3・Cl型で、地化学温度計に よる推定では温泉貯留層の温度は142-177℃と推定された.本研究で推定した温泉水の起源、水 質および貯留層温度はいずれも既往研究による報告と同質・同等であり、今回観察された枯渇 源泉からの高温泉の再湧出が最近の焼岳の火山活動の変化に起因する証拠は見出せなかった.

キーワード:焼岳,奥飛騨温泉郷,地震,異常湧出,化学・安定同位体(δD,δ¹⁸O)組成,地 化学温度計

Abstract

Since late April 2020, earthquakes in the central Nagano Prefecture and the Hida region of Gifu Prefecture have increased. During this seismic activity, an abnormal discharge of high-temperature hot spring water was observed at the abandoned well in the Okuhida hot spring area. As the Okuhida hot spring area is located about 4 km west of Yakedake volcano, the major chemical and stable isotopic compositions of hydrogen and oxygen of the discharged

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water were analyzed to investigate whether this abnormal discharge phenomena is caused by the volcanic activity of Yakedake volcano.

The analysis revealed the water was predominantly meteoric in origin, having a neutral pH and Na-HCO₃ \cdot Cl type, and the reservoir temperature was estimated to be 142–177°C by the geochemical thermometer. The origin, water quality, and reservoir temperature estimated in this study were comparable to those in previous studies ; no evidence was found that the abnormal discharge observed at the abandoned well was caused by recent changes in the volcanic activity of Yakedake volcano.

Key words : Yakedake volcano, Okuhida hot spring area, earthquakes, abnormal discharge, chemical and isotopic (δD and $\delta^{18}O$) compositions, geothermometer

1. Introduction

Since approximately 2:00 am on April 22, 2020, frequent earthquakes have occurred in the central Nagano Prefecture and the Hida region of Gifu Prefecture (Fig. 1a and b). In this region, seismicity was also active in 1998, and the maximum seismic intensity of 5-lower was observed on August 7, 1998 (Gifu Meteorological Office, 2020). In 2020 activity, the maximum seismic intensities of 4 were observed on April 23 in Matsumoto City of Nagano Prefecture and on May 19 and 29 in Takayama City of Gifu Prefecture, and after that, earthquakes with seismic intensity 2-1 continued to occur (Fig. 1c and Gifu Meteorological Office, 2020 a and b). The recent earthquake information is available on the web site of the seismic intensity database of Japan Meteorological Agency.

During this seismic activity, an abnormal discharge of high-temperature hot spring water from an abandoned hot spring source well in the Okuhida hot spring area of Takayama City was observed. The Okuhida hot spring area consists of five hot spring groups, Shin-Hodaka, Tochio, Shin-Hirayu, Fukuji, and Hirayu Onsen (Fig. 2). There are many high-temperature hot springs in this area (maximum outlet temperature is about 98° C), and one of the assumed heat sources of these hot springs is an active volcano, such as Yakedake that is located 4 km east of the hot spring area (Furuta, 2020). Based on chemical and isotopic analyses, the origin of those hot spring waters is considered mostly meteoric, but it has been pointed out that the hot spring waters may contain a small amount of volcanic fluid (Mizutani *et al.*, 1980 ; Koji *et al.*, 2016). Therefore, in this study, the major chemical and stable isotopic compositions of hydrogen and oxygen of the water sampled at the abandoned well were analyzed to investigate whether the abnormal discharge is caused by the volcanic activity.

2. Hot spring sample and Analyses

Abnormal discharge of high-temperature hot spring was observed at an abandoned well of the Okuhida Garden Hotel YAKEDAKE in the Shin-Hirayu hot spring group (Fig. 2). According to an interview with the hotel, detailed information on this abandoned well, such as a drilling data and hot spring analysis data, has already been lost, but the bottom depth of this well was 50 m, the outlet temperature was 70°C or more, and the discharged rate was about 500 L/min, at the time when the source was in use. This well stopped about 30 years ago due to scale clogging



Fig. 1 Seismic activity around the Gifu-Nagano district (modified after original data from Japan Meteorological Agency). (a) Index map. (b) Epicentral map. (c) Magnitude—Time diagram of earthquakes.

and was abandoned. The well was once cleaned more than 10 years after it stopped, but the hot spring water was not discharged and was abandoned.

The abnormal discharge reported in this paper was a temporary phenomenon, observed twice from 07:00 on May 10 to 15:00 on the following day and from 07:00 to 09:20 on May



Fig. 2 Index map of sampling site and Okuhida hot spring area. Base map was created by overlaying the standard and shaded relief map of the Geographical Survey Institute of Japan and a seamless geological map of the Geological Survey of Japan. The double circle (◎) shows the abandoned well. Cross marks (×) are source wells currently in use near the abandoned well. The solid lines show the fault and estimated fault.

19 (Fig. 3a and b). Afterward, the water discharge stopped again (Fig. 3c). The water was collected at approximately 08 : 30 on May 19, and the temperature was 77°C.

Analyses were conducted at the Meteorological Research Institute. The pH and electrical conductivity were measured using a handheld pH/EC meter (Horiba, D-74) equipped with a glass electrode (Horiba, 9615S-10D) for the pH measurement and a platinum-platinum black electrode (Horiba, 3552-10D) for the EC measurement. The concentrations of F, Cl, and SO₄ were measured by using ion chromatography (Thermo Scientific, Dionex Integrion HPIC), and Na, K, Ca, Mg SiO₂ were measured by using a microwave plasma atomic emission spectrometer (Agilent, 4210 MP-AES). The total alkalinity as HCO_3^- was determined by titration with H_2SO_4 to the final pH of 4.8 (Hach, AL-DT). Stable isotope ratios (δ D and δ^{18} O) were analyzed by using a cavity ring-down spectrometer (Picarro, L2140-i) and reported in per mil (‰) relative to Vienna Standard Mean Ocean Water (V-SMOW). The analytical results are listed in Table 1.



Fig. 3 Photograph of the abandoned well. During the abnormal discharge (May 19, 2020; a and b). After the discharge stops (May 27, 2020; c).

3. Results and Discussion

3.1 Isotopic composition

As the δD and $\delta^{18}O$ values of hot spring waters in the Okuhida area are almost the same as the value of river water in the area, the hot spring waters are considered basically meteoric. However, it is suspected that the hot spring waters in this area either contain volcanic water or are affected by evaporation because some hot spring waters have δD and $\delta^{18}O$ values close to the volcanic gases of Yakedake volcano (Sugiura and Mizutani, 1978; Mizutani *et al.*, 1980; Koji *et al.*, 2016). Further investigation is needed to determine whether the reason for such a heavy isotope ratio is due to the contribution of volcanic gas or evaporation, but in any case, the δD and $\delta^{18}O$ values of the hot spring water collected in this study is in good agreement with the meteoric line; there is no effect of volcanic water or evaporative concentration (Fig. 4).

3.2 Geochemical feature of the sampled water

In general, hot springs containing magmatic volatiles have low pH and are rich in $SO_4^{2^-}$ and Cl (e.g., Iwasaki, 1976). However, the hot spring water collected in this study has a neutral pH of 7.23 with a Na-HCO₃ · Cl type and does not contain much $SO_4^{2^-}$ (Table 1). This quality is very

Table 1 Analytical results of	the discharged w	ater sample	ed on May 19), 2020, witl	n reference v	alues for the (Okuhida he	ot spring ar	ea.
Hot springs	Sample Date	Temp	EC.	Hq	Ŀ	C	SO₄	Na	х
		ပံ	S/m		mg/L	mg/L	mg/L	mg/L	mg/L
Discharged water from abandoned well ^[1]	May 19, 2020	Ц	0.306	7.23	1.1	398	3.6	532	37.3
		Data below	/ the broken line	are reference	s				
Source well near the abandoned well 1 ^[2]	Nov. 16, 2000	57.2	N.D.	6.67	1.0	382	0.5	514	35.3
Source well near the abandoned well 1 ^[3]	July 7, 2010	70.8	N.D.	7.00	1.2	350	N.D.	527	38.5
Source well near the abandoned well 2 ^[4]	Sep. 26, 2011	67.1	N.D.	69.9	1.0	152	5.9	280	25.6
Source well near the abandoned well 3 ^[5]	July 5, 2013	56.3	N.D.	6.68	0.9	96	6.4	195	18.9
Shin-Hirayu ^[6]	May – Sep., 2012	28.7-68.2	0.037 - 0.34	5.75-7.21	N.D.	13-422	14 - 32	35 - 614	2.7-29
Hirayu ^[6]	May – Sep., 2012	49.0-74.8	0.121 - 0.0184	6.10 - 6.63	N.D.	86 – 281	34-222	121–258	7.6-26
Fukuji ^[6]	May, 2012	47.7-67.1	0.086 - 0.123	6.25 - 6.58	N.D.	13 - 90	15 - 33	116 - 193	12 - 19
Tochio ^[6]	Sep., 2012	63.2	0.08	7.14	N.D.	123	13	121	5.2
Shin-Hodaka ^[6]	May, 2012	63.0-82.3	0.068-0.174	6.27-7.74	N.D.	70-230	33-51	125-333	12–36
Hot springs	Ca	Mg	нсоз	SiO ₂	δD	δ ¹⁸ Ο	G	sothermomete	
	mg/L	mg/L	mg/L	mg/L	%	~ %	Na-K ^[7]	Na-K-Ca ^[8]	SiO _{2^[9]}
Discharged water from abandoned well ^[1]	71.0	26.0	1305	193	-81.6	-11.26	152	142	177
		Data below	/ the broken line	are reference	s				
Source well near the abandoned well 1 ^[2]	117	50.8	1408	135	N.D.	N.D.	150	124	155
Source well near the abandoned well 1 ^[3]	110	31.2	1407	135	N.D.	N.D.	156	130	155
Source well near the abandoned well 2 ^[4]	99.1	35.8	982.0	202	N.D.	N.D.	178	108	181
Source well near the abandoned well 3 ^[5]	85.5	30.8	828.6	180	N.D.	N.D.	184	96	173
Shin-Hirayu ^[6]	23-106	3.2 – 65	160 - 1183	61 - 143	-80.474.5	-11.7211.15	117-230	35–125	111–157
Hirayu ^[6]	39-274	10 - 92	151 - 780	52 - 146	-84.580.4	-12.58–-11.56	104 - 222	58 - 125	102 - 158
Fukuji ^[6]	43-53	3.3-16	465 545	122 - 130	-78.175.8	-11.46 11.40	188-197	91 - 108	147–151
Tochio ^[6]	32	0.2	208	119	-79.6	-11.87	112	69	146
Shin-Hodaka ^[6]	16 - 49	1.8-12	124 - 510	66-213	-83.676.3	-11.4810.11	145 - 198	101 - 150	114–183
* N.D., No data.; E.C., Electrical conductivity; Si is expr 1.1 Sec. 1.2 AAA	essed as SiO ₂ .								
[1] Sampleu at the point markeu with 🥥 in Fig. 2. And [2] Samoo wall located at the saint mode of with "Y 1"	aryzeu m ums stuuy. in Fin 2 Data from the wol	liido af tha Oliid	Vida Gardan Hatal V	VAKEDAKE					
[2] Source well located at the point marked with " $\times 1$ "	in Fig. 2. Data from Nagano	o Pharmaceutica	Association (2010)).					
[4] Source well located at the point marked with " $ imes 2$ "	in Fig. 2. Data from Nagano	o Pharmaceutica	Association (2011)						
[5] Source well located at the point marked with " $ imes 3^{"}$	in Fig. 2. Data from Naganc	o Pharmaceutica	Association (2013)						
[6] Data from Koji et al., (2016). [7] Truesdell (1976).	[8] Fournier and Truesdell	l (1973). [9] Fou	ırnier (1977).						

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similar to the hot spring water currently in use at the Okuhida Garden Hotel YAKEDAKE (Table 1). Furthermore, the neutral Na-HCO₃ \cdot Cl type has been typical in the Okuhida hot spring area for more than 40 years (Mizutani *et al.*, 1980 ; Kawamoto *et al.*, 1984 ; Koji *et al.*, 2016 ; Furuta, 2020). Therefore, the effect of abnormal volcanic activity cannot be detected based on the chemical composition of the hot spring water that discharged from the abandoned well.

3.3 Estimation of reservoir temperature using geothermometer

If the volcanic activity of Yakedake increases, underground temperatures may rise. To investigate this possibility, the reservoir temperature of the hot spring discharged from the abandoned well was estimated by applying a geochemical thermometer. There are various methods and application conditions for geochemical thermometers, but here, I used a Na-K geothermometer (Truesdell, 1976; Equation 1), a Na-K-Ca geothermometer (Fournier and Truesdell, 1973; Equation 2), and a SiO₂ geothermometer (Fournier, 1977; Equation 3), which have been applied in the analysis of geothermal resources in the Okuhida hot spring area (Koji *et al.*, 2016).

Na-K (Truesdell, 1976):

$$T(C) = \frac{855.6}{\log (Na/K) + 0.857} - 273.15$$
(1)
*Na and K in mg/L

Na-K-Ca (Fournier and Truesdell, 1973):

$$T(^{\circ}C) = \frac{1647}{\log (Na/K) + 4/3 [\log (Ca^{1/2}/Na) + 2.06] + 2.47} - 273.15$$
*Na, K and Ca in mg/L
(2)

$$SiO_{2}$$
, no stream loss (Fournier, 1977) :
 $T(C) = \frac{1309}{5.19 - \log SiO_2} - 273.15$ (3)
 $*SiO_2$ in mg/L

As a result, the reservoir temperature of the hot spring water discharged from the abandoned well was estimated to be 142–177°C (Table 1). Although we must be cautious that the same source should be compared to investigate the temporal change of the reservoir temperature because the reservoir temperature varies even in a narrow area, the estimated temperature of 142–177°C is within the range estimated from the analytical value of the source wells near the abandoned well (96–184°C : Table 1) and is also reasonable as the previously estimated reservoir temperature in the Okuhida hot spring area (approximately 150–200°C : Mizutani *et al.*, 1980 ; Koji *et al.*, 2016). Therefore, based on this estimation, there is no evidence that the abnormal discharge observed at the abandoned well was caused by the underground temperature change due to the volcanic activity of Yakedake volcano.

4. Concluding remarks

During the seismic activity at the Gifu-Nagano district, abnormal discharge phenomena of high-temperature hot spring water from the abandoned well in the Okuhida hot spring area was observed. To investigate whether the abnormal discharge was caused by the volcanic activity of Yakedake volcano, the major chemical and stable isotopic compositions of hydrogen and oxygen were analyzed, and the reservoir temperature was estimated by applying geothermometers. As a result, no evidence was found that this abnormal discharge was caused by an increase of volcanic activity of Yakedake volcano.

Although the relationship between earthquakes and hot spring anomalies has not been completely clarified, many abnormal phenomena associated with earthquakes have been reported in Japan as well as overseas, such as changes in water level, discharge rate, outlet temperature, chemical composition, and color appearance (e.g., Nishimura, 1999; Okusawa and Tsukahara, 2002; Itaba and Koizumi, 2007; Itadera *et al.*, 2011; Sasaki, 2012; Cox *et al.*, 2015). As shown in Figure 2, the abandoned well, in which the abnormal discharge was observed this time, is located at the position where many faults exist, and it is assumed that the groundwater flow, including the hot spring water, in this area is restricted by these faults (e.g., Akiyama, 1978). Therefore, it may be reasonable to interpret that the abnormal discharge of the high-temperature hot spring water observed this time is caused by the change in the underground flow path due to the recent frequent earthquakes in the Gifu-Nagano district.

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